

Particulate Gravity Currents with Resuspension

Jim McElwaine

Professor of Geohazards
Department of
Earth Sciences
Durham University



Acknowledgments



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Kouichi Nishimura
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Plan of Talk

- Transition Experiments
- Direct Numerical Simulations
- Dynamic Models



Powder Avalanche on K2

Pierre Beghin



Head of Powder Snow Avalanche

Cemagref



Slab Avalanche Fracture Line

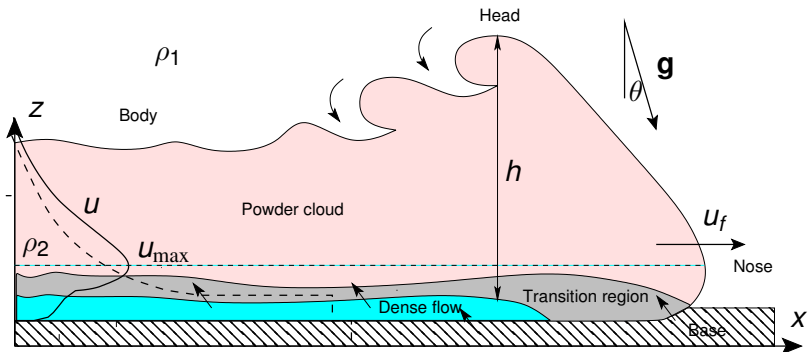


Test Chute in Davos

film



Schematic of Avalanche, Turbidity Current, Pyroclastic Flow



Aim

Understand formation of Suspension Currents

Use steep slopes to give a low Richardson number for large density difference

- Transition to suspension
- Limited entrainment
- Steady flows
- Understand air interaction
- Comparison of field observations with experiments

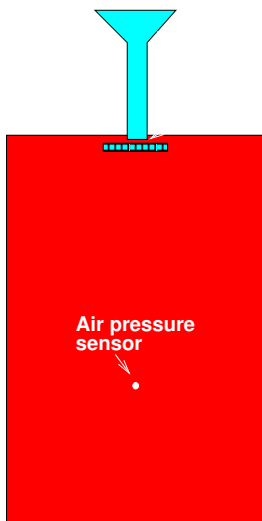
Similarity Criteria

Experiment	Materials	Re_p	Ri	St	$\frac{\Delta\rho}{\rho_a}$	Re
Powder snow avalanches	Snow-air	3000	1	0.02	10	10^9
Ancey (2004)	sawdust water	50	1.7	0.006	0.05	10^4
Bozhinskiy (1998)	aluminium air	0.1	20	0.03	1	10^3
Beghin (1981) Beghin (1983) Beghin (1983)	Brine-water/ Sand-water suspension	–	5	–	0.02	10^4
Nishimura (1998) McElwaine (2001)	Ping-pong balls air	2×10^4	2	10	50	10^7
Hampton (1972)	Kaolinite and water slurry	–	< 0.5	–	0.1	–
Hermann (1987)	Polystyrene powder water	1.5	0.1	10^{-4}	0.002	10^4
Hopfinger (1977) Tochon-Danguy (1974)	Brine water	–	2	–	0.01	10^3
Present Study	Snow-air	150	1	10	10	10^5
Present Study	Polystyrene-air	150	2	1	5	10^4

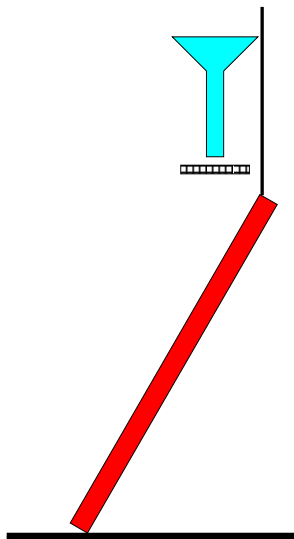
Similarity Summary

- Exact similarity of Ri and $\Delta\rho/\rho_a$
- Re and Re_p not matched but qualitatively similar
- St is different \Rightarrow sedimentation is important
- Slope angles are different.
Appears unimportant at high Re

Experiments



front view



side view

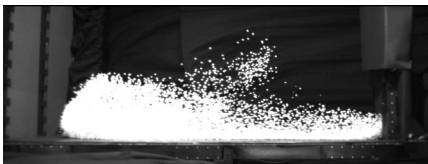
Side View 8 Litre Avalanches



31.5° slope



58.5° slope



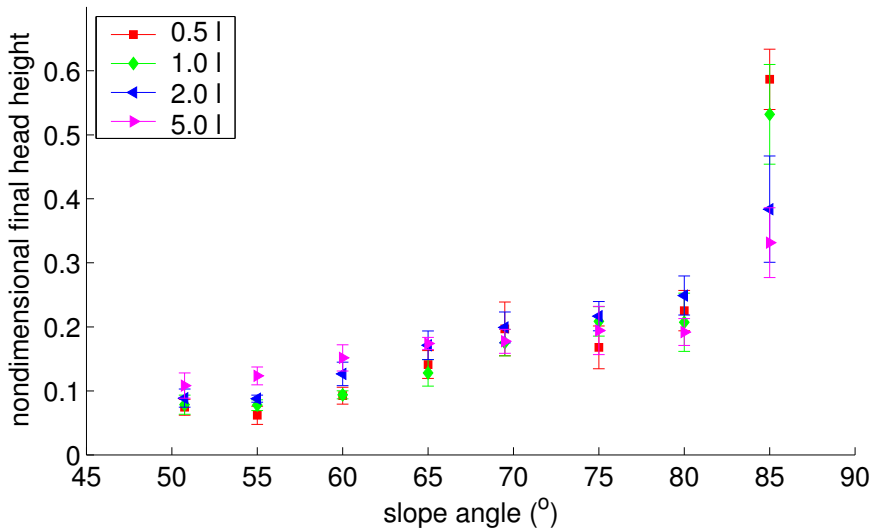
91.0° slope

100 ml side
8000 ml side

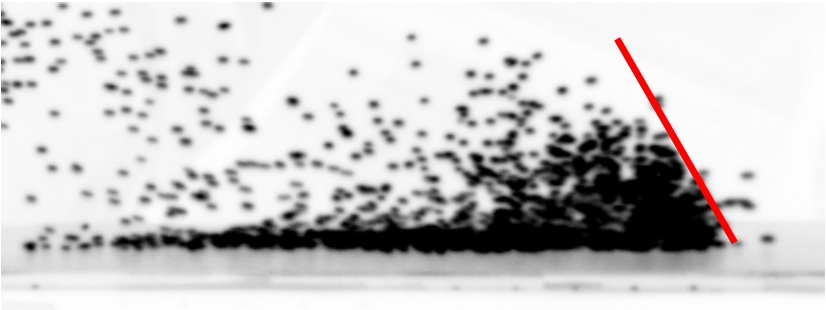
100 ml front
8000 ml front

Non-dimensional Height

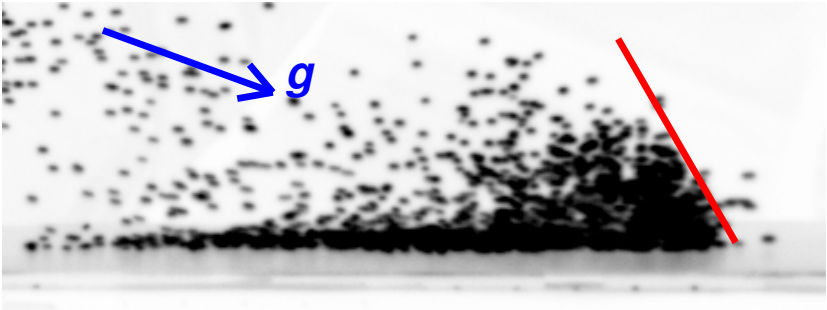
$$\tilde{h} = \frac{h}{V^{1/3}}$$



Polystyrene balls



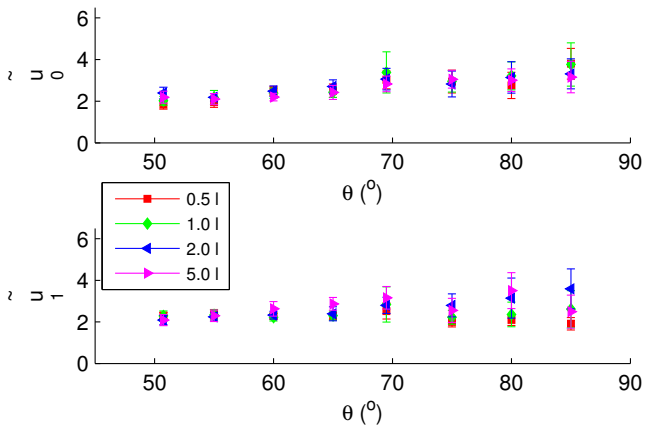
Polystyrene balls on 70° surface



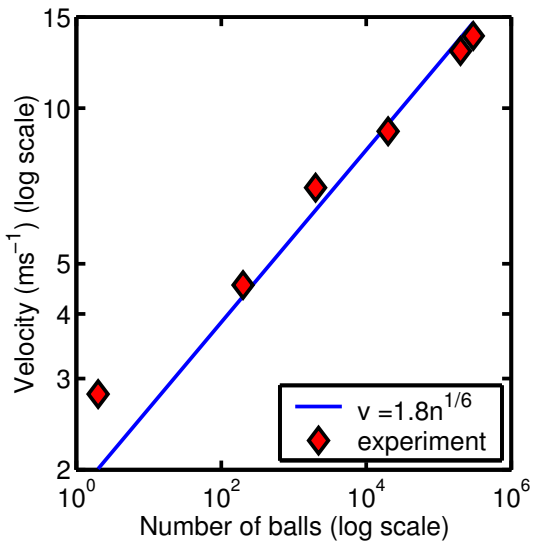
Ping-Pong Avalanches



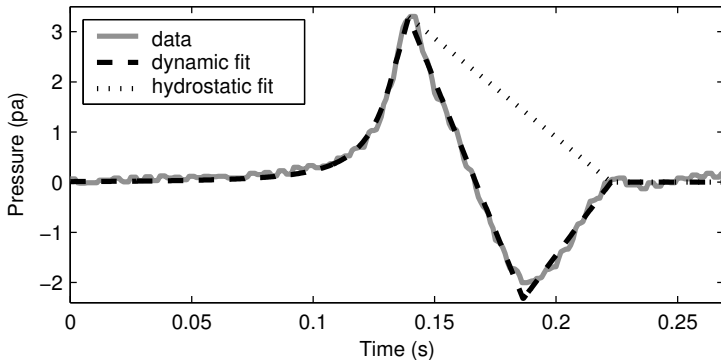
Non-Dimensional Velocity

$$\tilde{u} = \frac{u}{V_{\frac{1}{6}} g^{\frac{1}{2}}}$$


Front Velocities at the K-Point

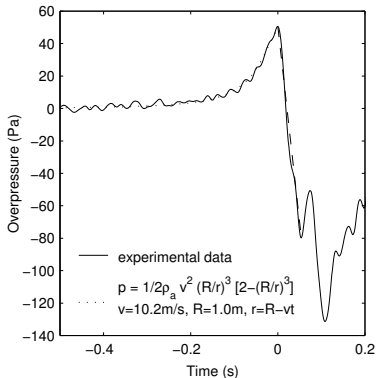


Pressure Theory Comparison

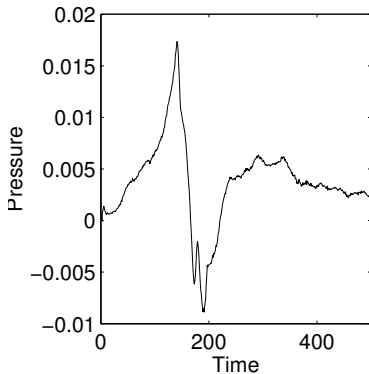


Air pressure in a small powder cloud

DNS simulations and Ping-Pong experiments



Air pressure data from a ping-pong ball avalanches and comparison with theory



Air pressure data from a direct numerical simulation of a gravity current

Front Instability, 51 polystyrene, 51 ° slope



Conclusions

- Transition to suspension can be achieved in the laboratory
- Can deduce avalanche **length**, **height**, **speed**, and **front angle** from pressure data
- Good agreement between theory, experiments and field observations
- Pressure measurements can distinguish suspended from dense flows
- Coherent internal velocities can be twice front velocity
Take care estimating forces !

Direct Numerical Simulations

- Meiburg Code
2d spectral with compact finite differences
- Diablo from John Taylor
3d spectral with low order finite differences
- Simulation region 8×1
- Release area 2×0.5
- Slope angles $0-90^\circ$
- Boussinesq and non-Boussinesq

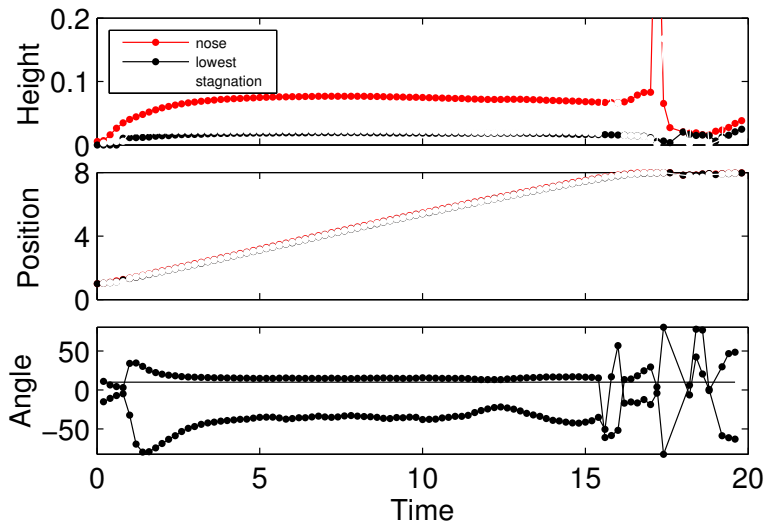
Test hypothesis:

stagnation point is lowest point as $Re \rightarrow \infty$

Time evolution, $Re=32,000$, $Slope=10^\circ$

film

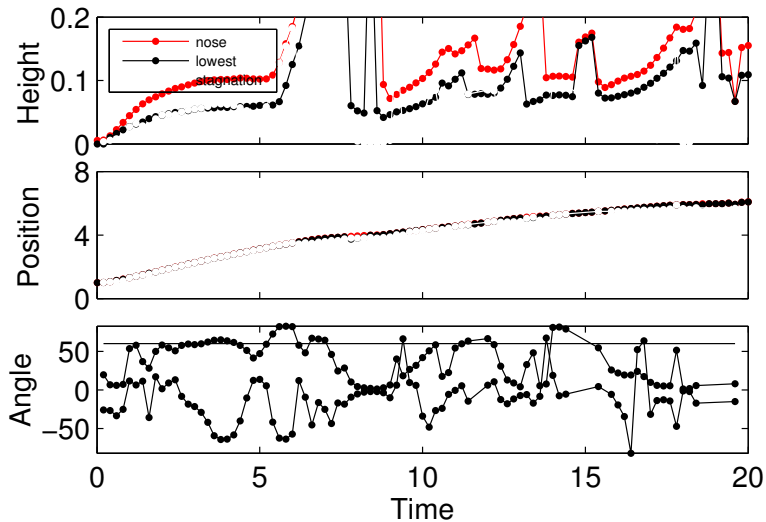
front



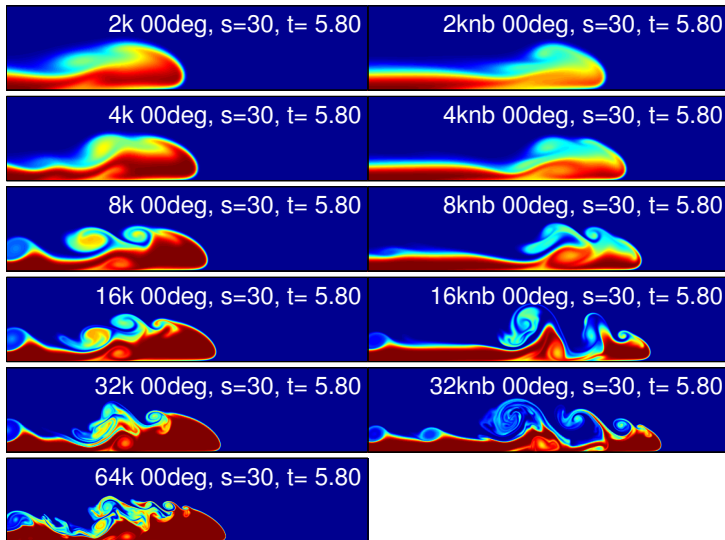
Time evolution, $Re=32,000$, $Slope=60^\circ$

film

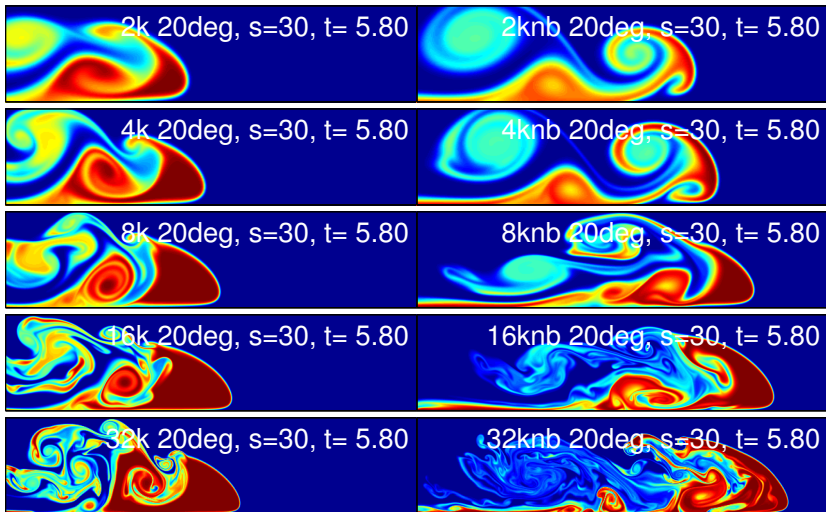
front



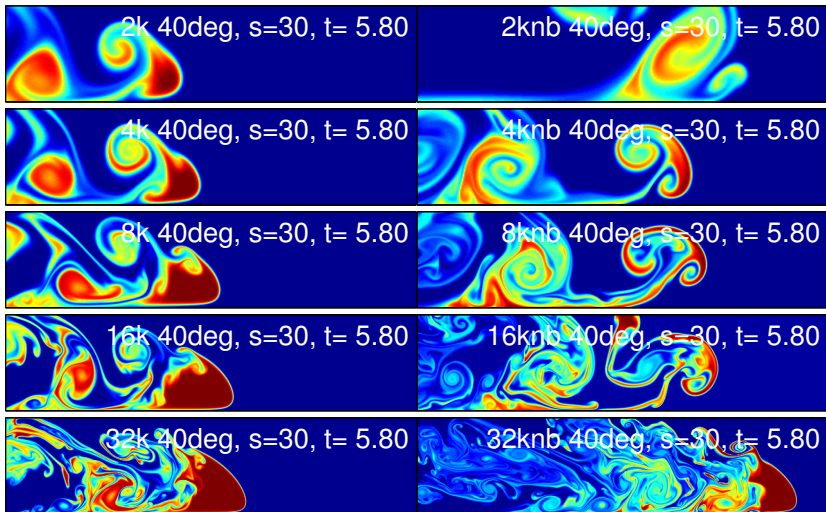
Re Comparison at slope 0°



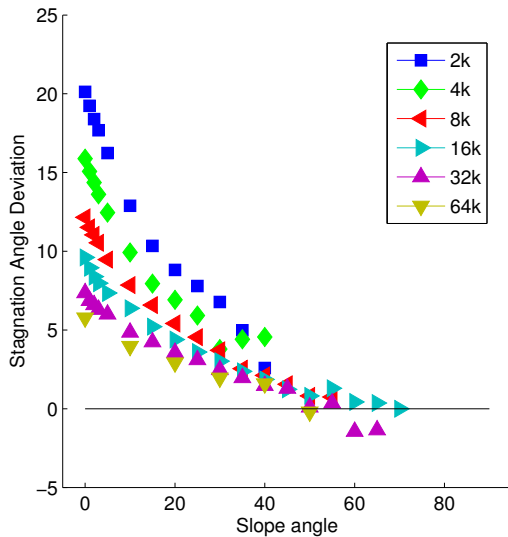
Re Comparison at slope 20°



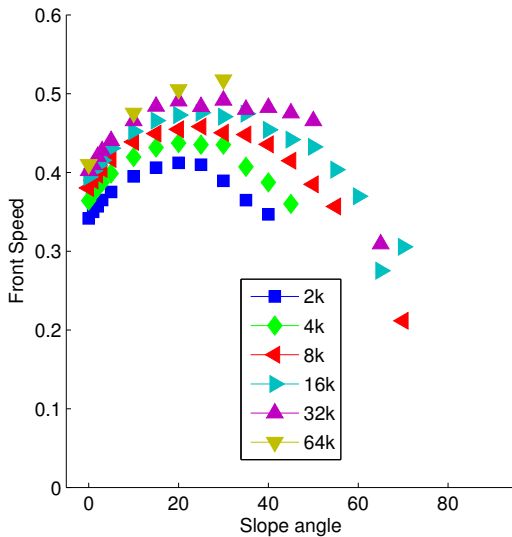
Re Comparison at slope 40°



Stagnation Point Angle



Front Speed — High angles evaporate



Hindered Sedimentation, with Particle Pressure

$$\frac{\partial \phi}{\partial t} + \nabla \cdot \mathbf{q} = 0,$$

where particle flux

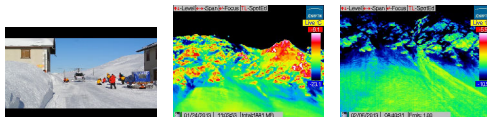
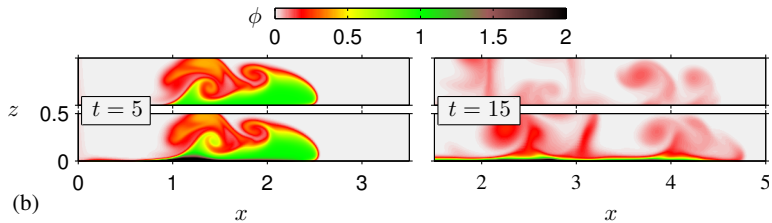
$$\mathbf{q} = \mathbf{u}\phi + \mathbf{u}_s\phi(1 - \alpha\phi) - \frac{D\nabla\phi}{1 - \beta\phi}$$

Sedimenting Boundary condition $\mathbf{n} \cdot \nabla(\mathbf{n} \cdot \mathbf{q}) = 0$ or $\nabla\phi = 0$

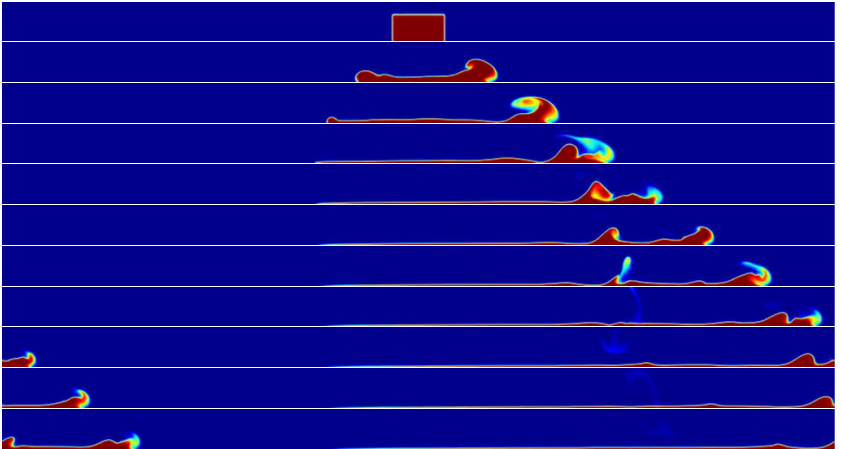
Resuspending boundary condition $\mathbf{n} \cdot \mathbf{q} = 0$

Use mixed compact finite differences and finite volume schemes to exactly conserve mass.

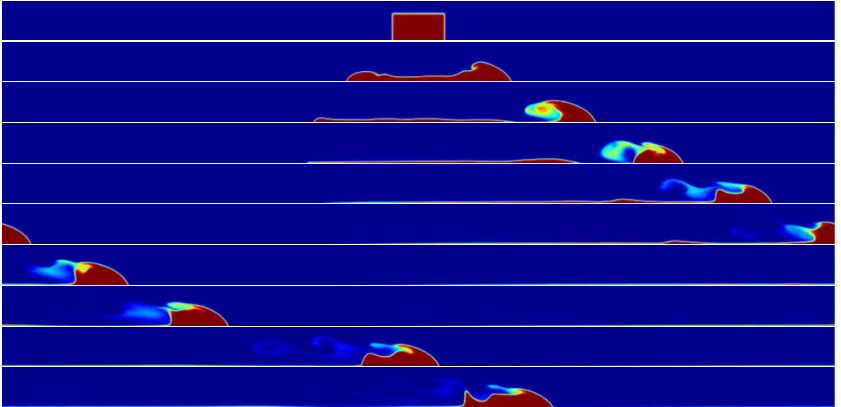
Sedimentation vs Resuspension



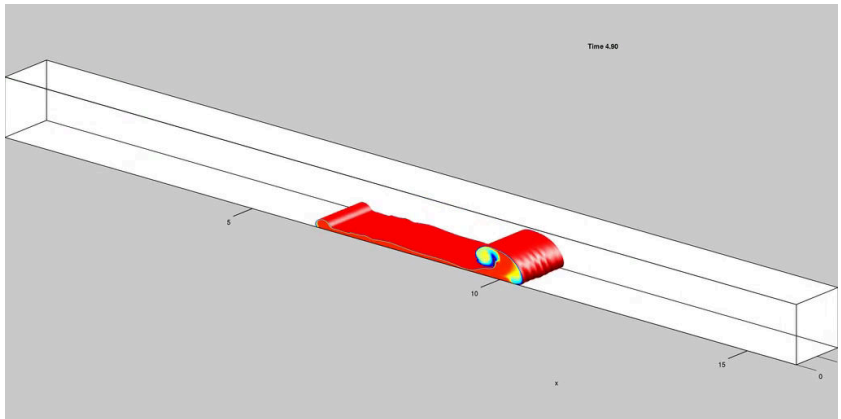
2D with no-slip and hindered sedimentation, $Re = 2000$



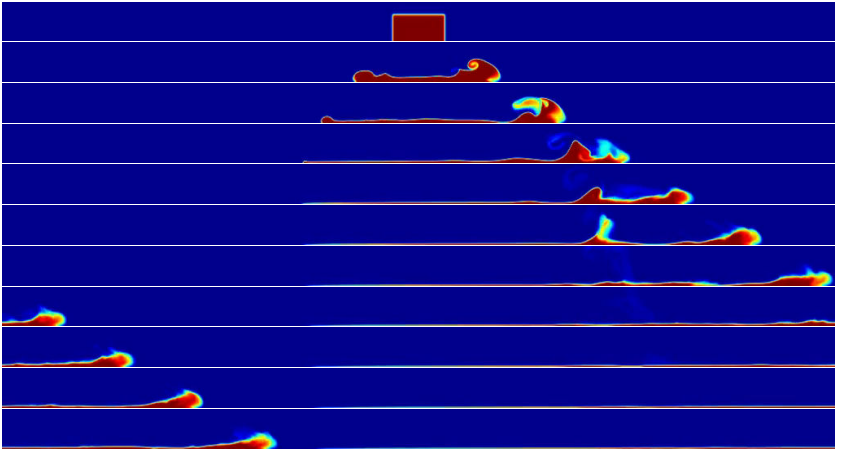
2D with slip and hindered sedimentation, $Re = 2000$



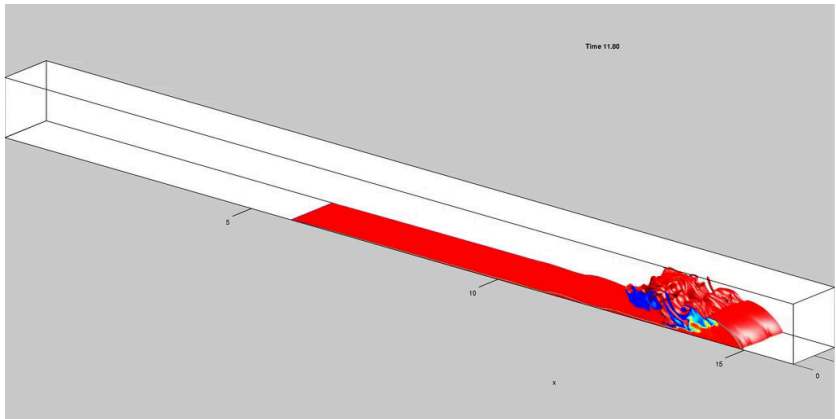
3D with no-slip and hindered sedimentation, $Re = 4\,000$



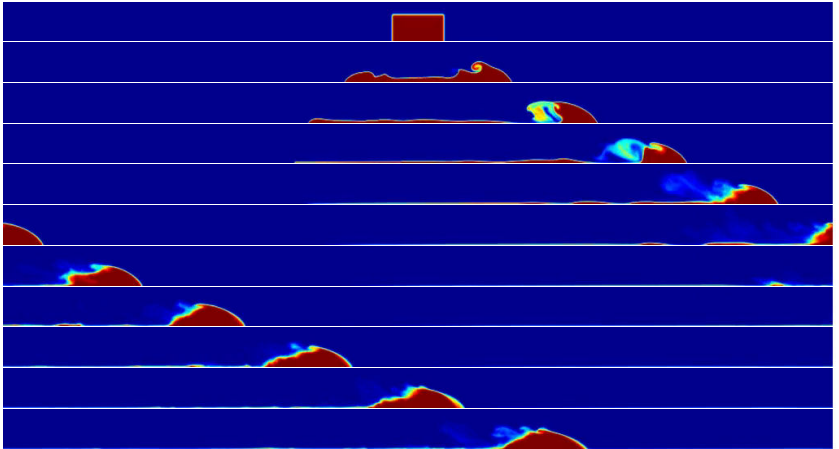
3D with no-slip and hindered sedimentation, $Re = 4\,000$



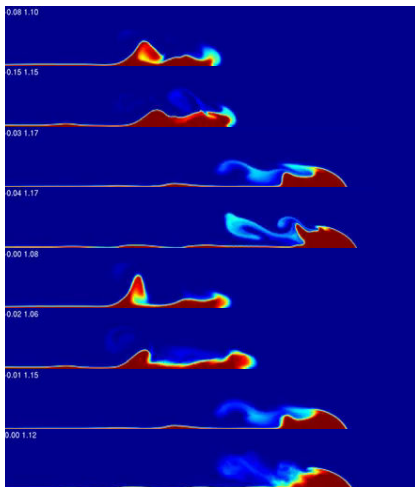
3D with slip and hindered sedimentation, $Re = 4\,000$



3D with slip and hindered sedimentation, $Re = 4\,000$



Comparison



2D Re=2 000 no-slip

2D Re=4 000 no-slip

2D Re=2 000 slip

2D Re=4 000 slip

3D Re=2 000 no-slip

3D Re=4 000 no-slip

3D Re=2 000 slip

3D Re=4 000 slip

Kulikovskiy–Sveshnikova–Beghin (KSB)

Three conservation equations

volume $\frac{dV}{dt} = q_s + q_a$

buoyancy $\frac{dB}{dt} = (\rho_s - \rho_a)q_s$

momentum $\frac{d}{dt} \left\{ \left[B + (1 + \chi) V \rho_a \right] u \right\} = Bg \sin \theta$

ρ_s snow density

g gravity

q_s snow flux

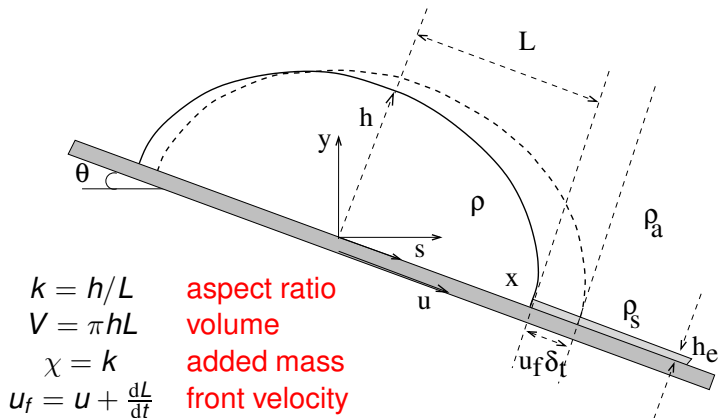
ρ_a air density

θ slope angle

q_a air flux

χ added mass

Geometric Closure



Flux Closures

$$q_s = u_f h_e - \beta u \sqrt{V} \quad \text{snow entrainment/detrainment}$$

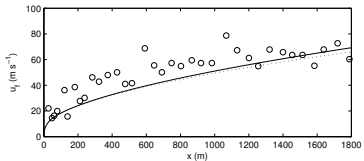
$$q_a = (\alpha u - u_s) \sqrt{V} \quad \text{air entrainment/detrainment}$$

$$\alpha(\text{Ri}) = \begin{cases} e^{-\lambda \text{Ri}^2} & \text{Ri} \leq 1 \\ \frac{e^{-\lambda}}{\text{Ri}} & \text{Ri} > 1 \end{cases}$$

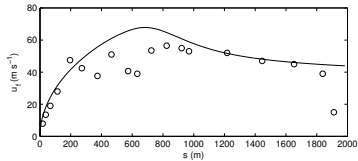
$$\text{Ri} = \frac{\rho - \rho_a}{\rho_a} \frac{gh \cos \theta}{u^2}$$

h_e erodible snow depth β mass loss coefficient
 u_s sedimentation velocity

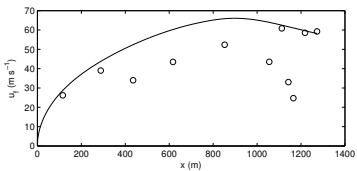
Comparison with Velocity Data



Avalanche no. 200

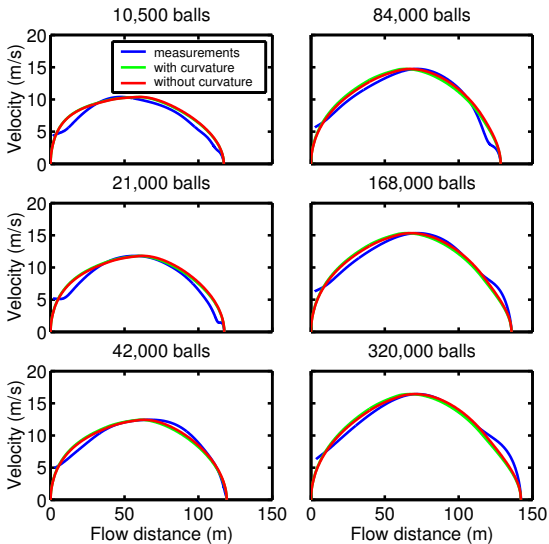


Avalanche no. 509

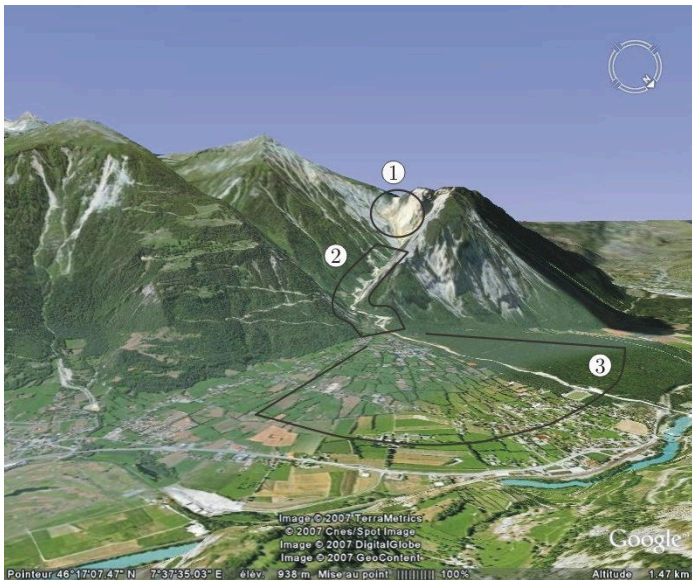


Avalanche no. 628

Comparison with Ping-Pong Avalanches



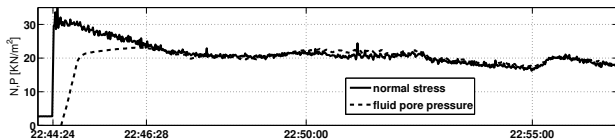
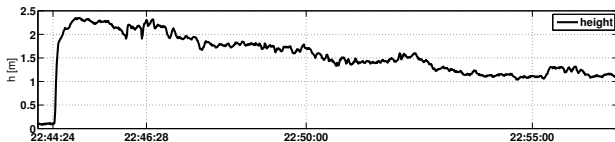
Illgraben Situation



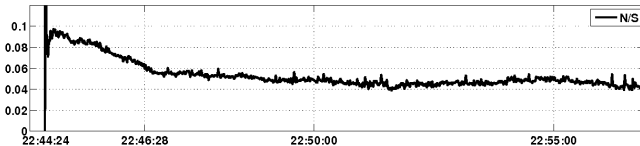
Illgraben Bridge



Height and Stress Data



Ratio of shear force to normal force

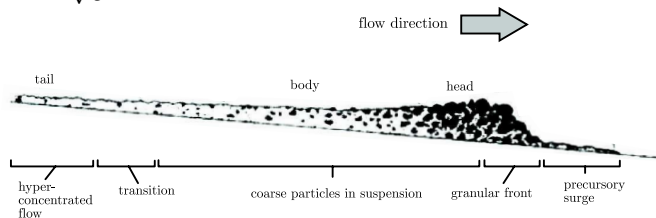


Motivation – Zeroth order shallow water

$$\frac{\partial u}{\partial t} = g \sin \theta - \mu g \cos \theta$$

- Chezy $\mu(\text{Fr}, h/d, \theta) \propto \text{Fr}^2$
- Viscous $\mu(\text{Fr}, h/d, \theta) \propto \text{Fr}$
- μ should also depend on solids concentration at the bottom

$$\text{Fr} = \frac{u}{\sqrt{gh \cos \theta}}$$



Two phase debris flow model

After much algebra . . .

vertical equation of mass conservation is

$$\partial_t \alpha = \partial_y [V\alpha(1 - \alpha) + D\partial_y \alpha] = \partial_y D [Pe \alpha(1 - \alpha) + \partial_y \alpha]$$

α density of solids relative to max

V sedimentation velocity

D diffusion

$Pe = V/D$ Peclet number

This is a diffusion equation with hindered settling and the same as Gray's segregation theory

Reduction to slow manifold

Define the vertical centre of mass for the solids fraction

$$h_p(t) = \int_0^H y \alpha(t, y) dy.$$

Then use previous result and more algebra to get

$$\frac{dh_p(t)}{dt} = \frac{h_p^* - h_p(t)}{T} [1 + \epsilon(h_p^* - h_p(t) + \dots)].$$

For resuspension dominated regime $h_p^* = \frac{H}{2}$ and mixture becomes well mixed

For sedimentation dominated regime $h_p^* = \frac{H_p}{2}$ where H_p is total height of particles at maximum packing fraction

Excess pore pressure is

$$\frac{p}{\rho_f g_y} = \frac{c_1(2h_p - H_p) + c_2(h_p - h_p^*)}{H - H_p}$$

Complete Set of Depth Averaged Equations

$$\partial_t M + \partial_x (MU) = 0$$

$$\partial_t (cM) + \partial_x [cMU(1 + a_1 m_1)] = 0$$

$$\partial_t (UM) + \partial_x [MU^2(a_2 + a_3 m_1^2) + g_y hM] = Mg_x - \mu Mg_y$$

$$\partial_t m_1 + \partial_x [m_1 U + (a_4 + a_5 m_1^2)U] = \frac{m_1^* - m_1}{T_m} [1 + \dots]$$

$m_1 = 2h_p/H - 1$ dimensionless deviation from mixed

M total mass hold up

c relative concentration of particles

h centre of mass

Conclusions

- Resuspension very important in many geophysical flows
- DNS expensive but can reproduce two layer structure
- Integral point mass models work well for some cases
- Depth integrated equations with vertical redistribution a good compromise
- Easy to construct empirical models
- Also works for n component mixtures and segregation

Publications

- N.A. Konopliv and S.G. Llewellyn-Smith and **J.N. McElwaine** and E. Meiburg. Modeling gravity currents without an energy closure. *J. Fluid Mech.*, **789**:806–829, 2016.
- B. Sovilla, **J.N. McElwaine**, and M.Y. Louge. The structure of powder snow avalanches. *Comptes Rendus Physique*, **16**(1):97–104, 2015.
- B. Turnbull, E.T. Bowman, and **J.N. McElwaine**. Debris flows: Experiments and modelling. *Comptes Rendus Physique*, **16**(1):86–96, 2015.
- M. Ash, P.V. Brennan, C.J. Keylock, N.M. Vriend, **J.N. McElwaine**, and B. Sovilla. Two-dimensional radar imaging of flowing avalanches. *Cold Regions Science and Technology*, **102**:41–51, 2014.
- A.J. Hafiz, **J.N. McElwaine**, and C.P. Caulfield. The instantaneous Froude number and depth of unsteady gravity currents. *J. Hydraulic Res.*, **51**:432–445, 2013.
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- N. M. Vriend, **J.N. McElwaine**, B. Sovilla, C. J. Keylock, M. Ash, and P. V. Brennan. High resolution radar measurements of snow avalanches. *Geo. Res. Let.*, **40**(4):727–731, 2013.
- B. Turnbull and **J.N. McElwaine**, 2008. Experiments on the non-Boussinesq Flow of Self-Igniting Suspension Currents on a Steep Open Slope., *J. Geophys. Res.*, **113**(F01003).
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- Turnbull, B., and **J.N. McElwaine**, 2007. A Comparison of Powder Snow Avalanches at Vallée de la Sionne with Plume Theories, *J. Glaciol.*, **53**(30)
- **J.N. McElwaine**, and B. Turnbull, 2006. Plume Theories Versus Compact Models for Powder Snow Avalanches, *Sixth International Symposium on Stratified Flows, Perth, December 11-14*,
- **J.N. McElwaine**, 2005. Rotational flow in gravity current heads, *Phil. Trans. R. Soc. Lond.*, **363**, 1603–1623, 10.1098/rsta.2005.1597.
- **J.N. McElwaine** and B. Turnbull, 2005. Air Pressure Data from the Vallée de la Sionne Avalanches of 2004, *J. Geophys. Res.*, **110**(F03010).
- **J.N. McElwaine** and K. Nishimura, 2001. Particulate Gravity Currents, Blackwell Science, chap. Ping-pong Ball Avalanche Experiments, no. 31 in Special Publication of the International Association of Sedimentologists, 135–148.